

# CONSIDER A SPHERICAL EARTH HEAT CAPACITY, TIME CONSTANT AND SENSITIVITY OF EARTH'S CLIMATE SYSTEM

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<http://www.ecd.bnl.gov/steve>

# Consider a Spherical Cow

A Course in Environmental Problem Solving

JOHN HARTE



longer, directional force  
resisting downward motion

mg = symmetrical force downward



$2.469 \times 10^{21}$   $m = -21 - 222$



# OVERVIEW

Earth 's energy balance and perturbations

Climate sensitivity – definition, importance, past and current estimates

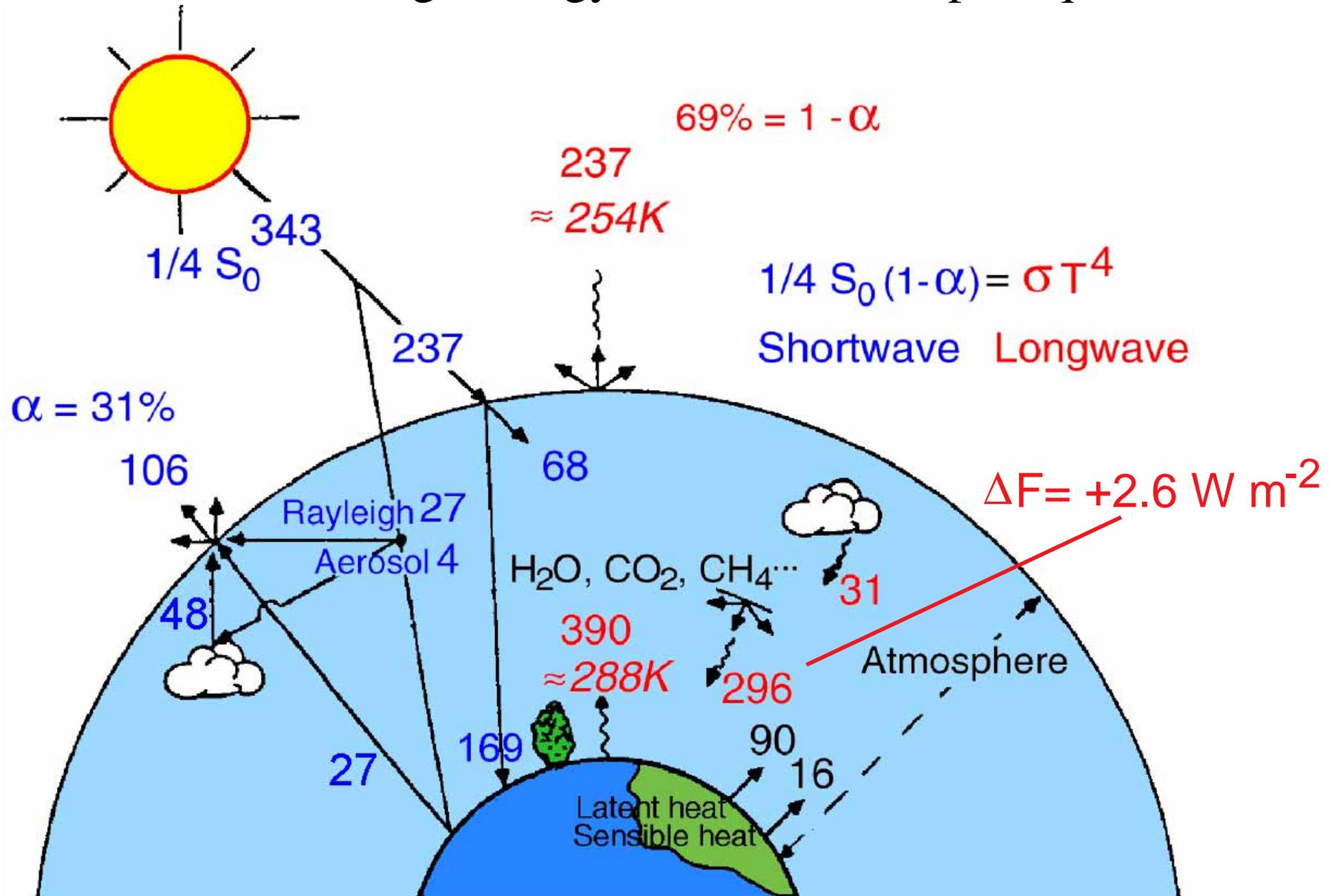
Previous approaches to Earth's climate sensitivity

Climate sensitivity from whole-earth energy-balance model

Concluding remarks

# GLOBAL ENERGY BALANCE

Global and annual average energy fluxes in watts per square meter



Schwartz, 1996, modified from Ramanathan, 1987

# ***ATMOSPHERIC RADIATION***

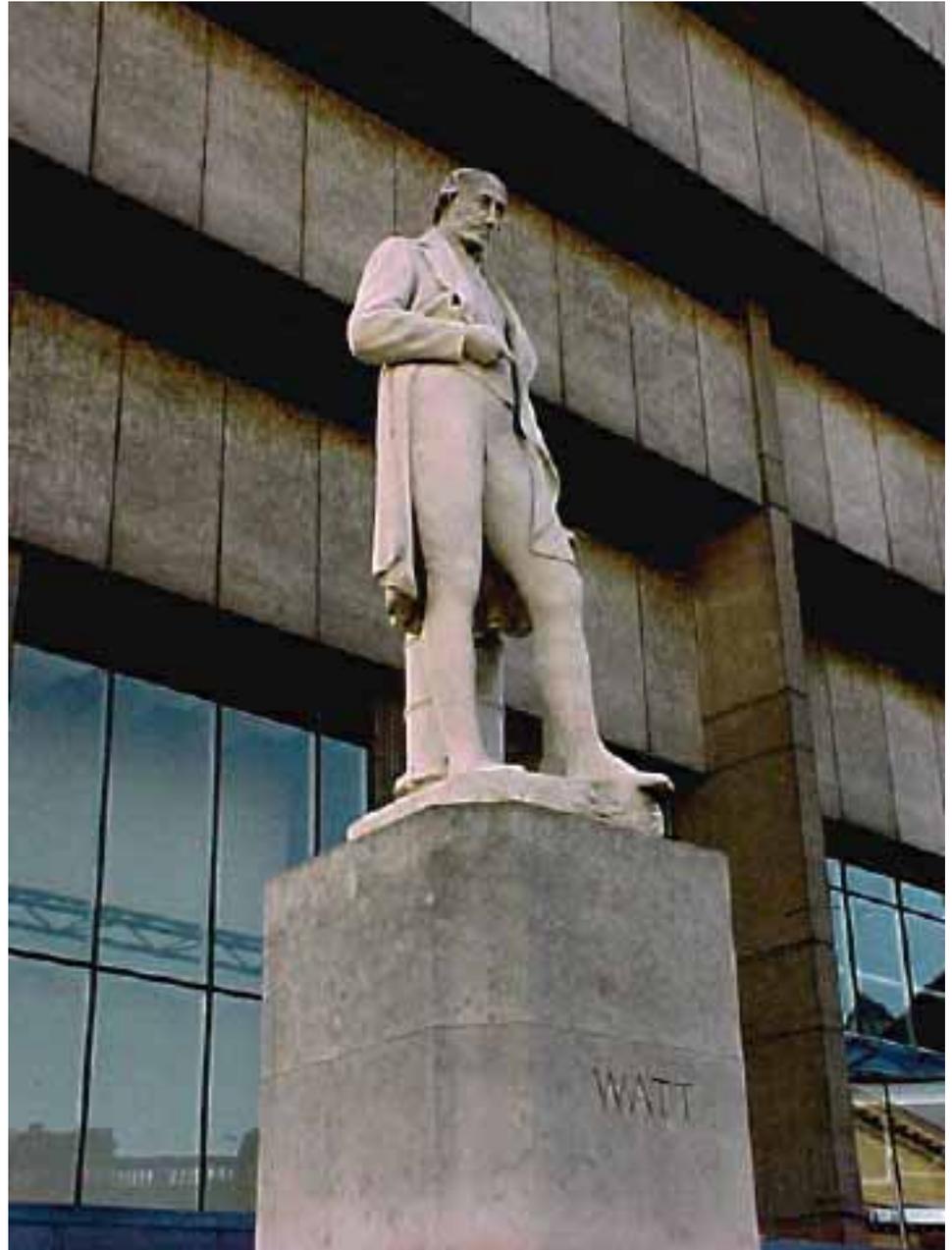
***Power per area***

***Energy per time per  
area***

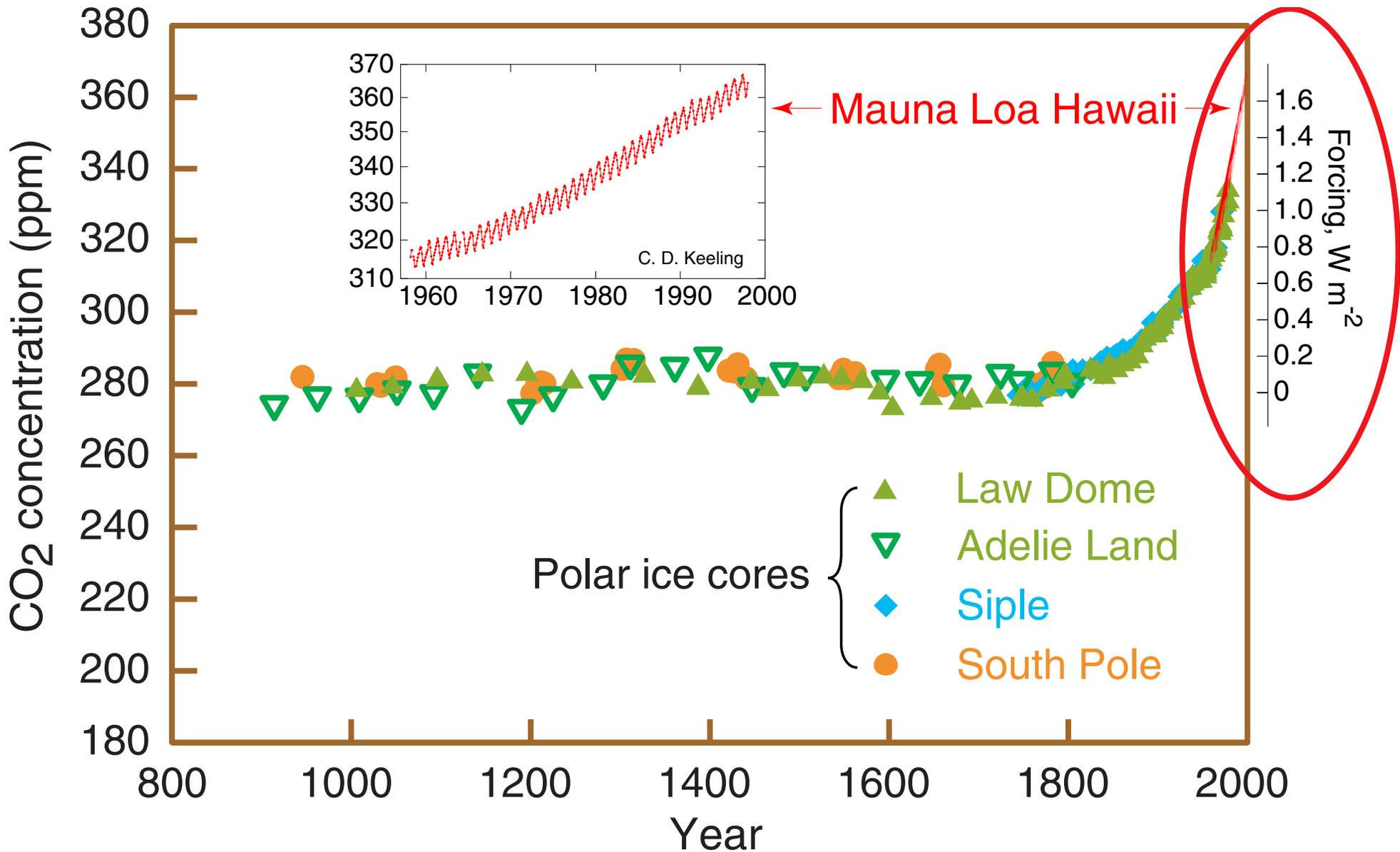
***Unit:***

***Watt per square meter***

***$W m^{-2}$***



# ATMOSPHERIC CARBON DIOXIDE IS INCREASING



Global carbon dioxide concentration and infrared radiative forcing over the last thousand years

# ***RADIATIVE FORCING***

A *change* in a radiative flux term in Earth's radiation budget,  $\Delta F$ ,  $\text{W m}^{-2}$ .

## ***Working hypothesis:***

*On a global basis radiative forcings are additive and fungible.*

- This hypothesis is fundamental to the radiative forcing concept.
- This hypothesis underlies much of the assessment of climate change over the industrial period.

# ***CLIMATE RESPONSE***

The *change* in global and annual mean temperature,  $\Delta T$ , K, resulting from a given radiative forcing.

***Working hypothesis:***

*The change in global mean temperature is proportional to the forcing, but independent of its nature and spatial distribution.*

$$\Delta T = S \Delta F$$

# *CLIMATE SENSITIVITY*

The *change* in global and annual mean temperature per unit forcing,  $S$ ,  $\text{K}/(\text{W m}^{-2})$ ,

$$S = \Delta T / \Delta F.$$

Climate sensitivity is not known and is the objective of much current research on climate change.

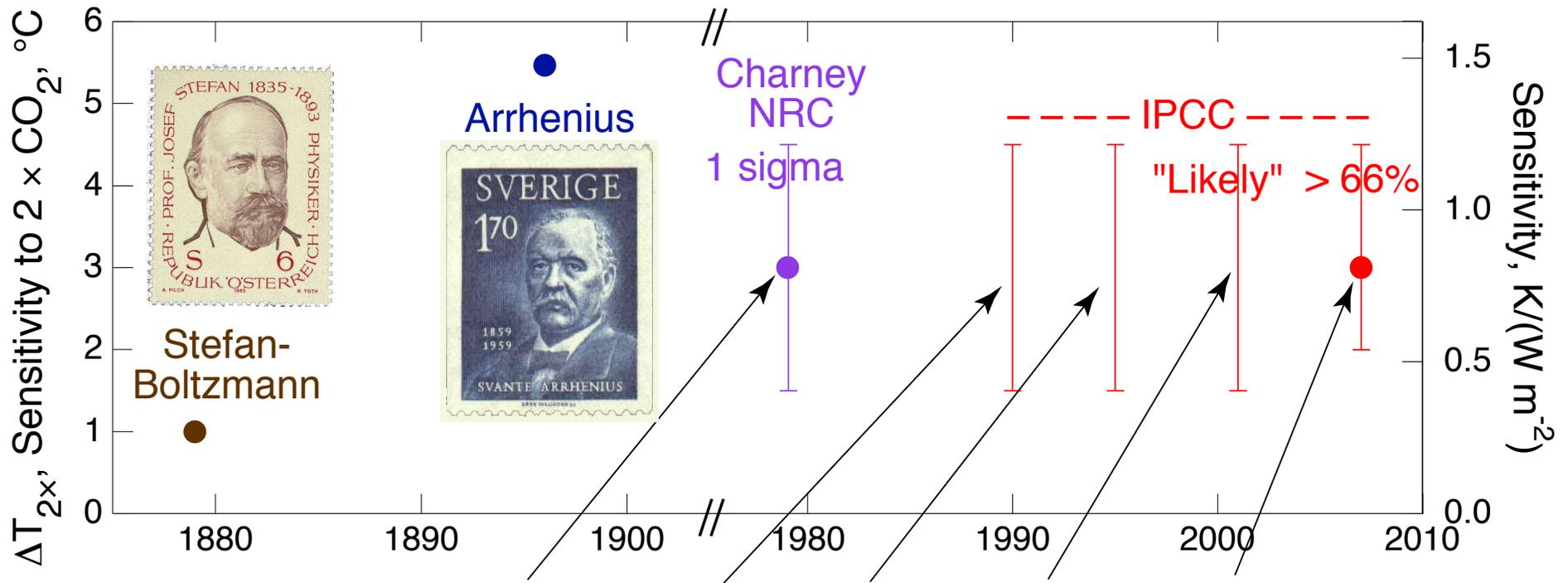
Climate sensitivity is often expressed as the temperature for doubled  $\text{CO}_2$  concentration  $\Delta T_{2\times}$ .

$$\Delta T_{2\times} = S \Delta F_{2\times}$$

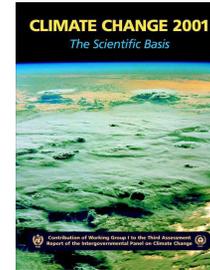
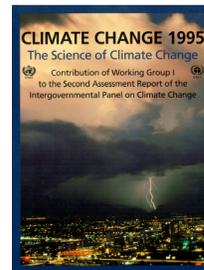
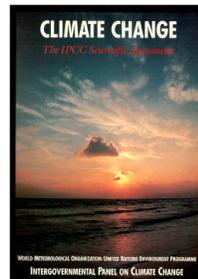
$$\Delta F_{2\times} \approx 3.7 \text{ W m}^{-2}$$

# CLIMATE SENSITIVITY ESTIMATES THROUGH THE AGES

Estimates of central value and uncertainty range from major national and international assessments



**Carbon Dioxide and Climate:  
A Scientific Assessment**  
NATIONAL ACADEMY OF SCIENCES  
Washington, D.C. 1979



Despite extensive research, climate sensitivity remains *highly uncertain*.

# ***IMPLICATIONS OF UNCERTAINTY IN CLIMATE SENSITIVITY***

Uncertainty in climate sensitivity translates directly into . . .

- Uncertainty in the amount of *incremental atmospheric CO<sub>2</sub>* that would result in a given increase in global mean surface temperature.
- Uncertainty in the amount of *fossil fuel carbon* that can be combusted consonant with a given climate effect.

***At present this uncertainty is about a factor of 3.***

# ***IMPORTANCE OF KNOWLEDGE OF CLIMATE TO INFORMED DECISION MAKING***

- The lifetime of incremental atmospheric CO<sub>2</sub> is about 100 years.
- The expected life of a new coal-fired power plant is 50 to 75 years.

***Actions taken today will have long-lasting effects.***

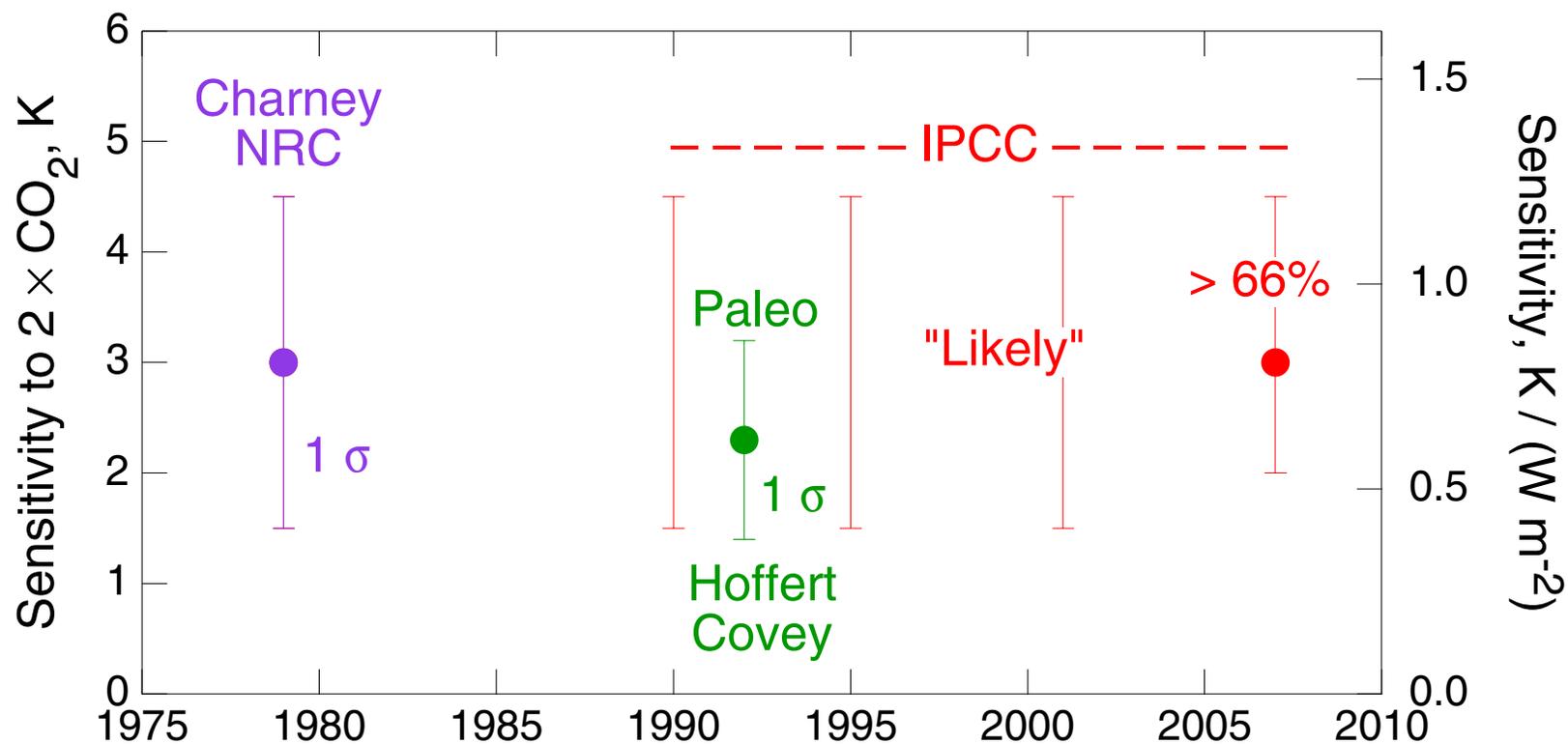
***Early knowledge of climate sensitivity can result in huge averted costs.***

# ***KEY APPROACHES TO DETERMINING CLIMATE SENSITIVITY***

- *Paleoclimate studies*: Forcing and response over time scales from millennial to millions of years.

# CLIMATE SENSITIVITY ESTIMATES THROUGH THE AGES

Estimates of central value and uncertainty range from specific approaches and major national and international assessments



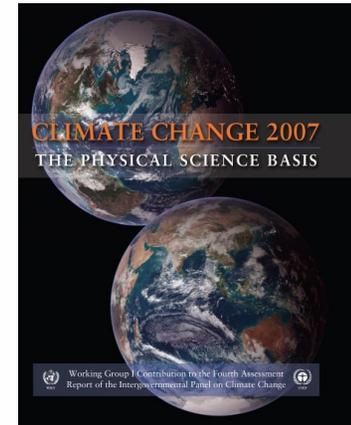
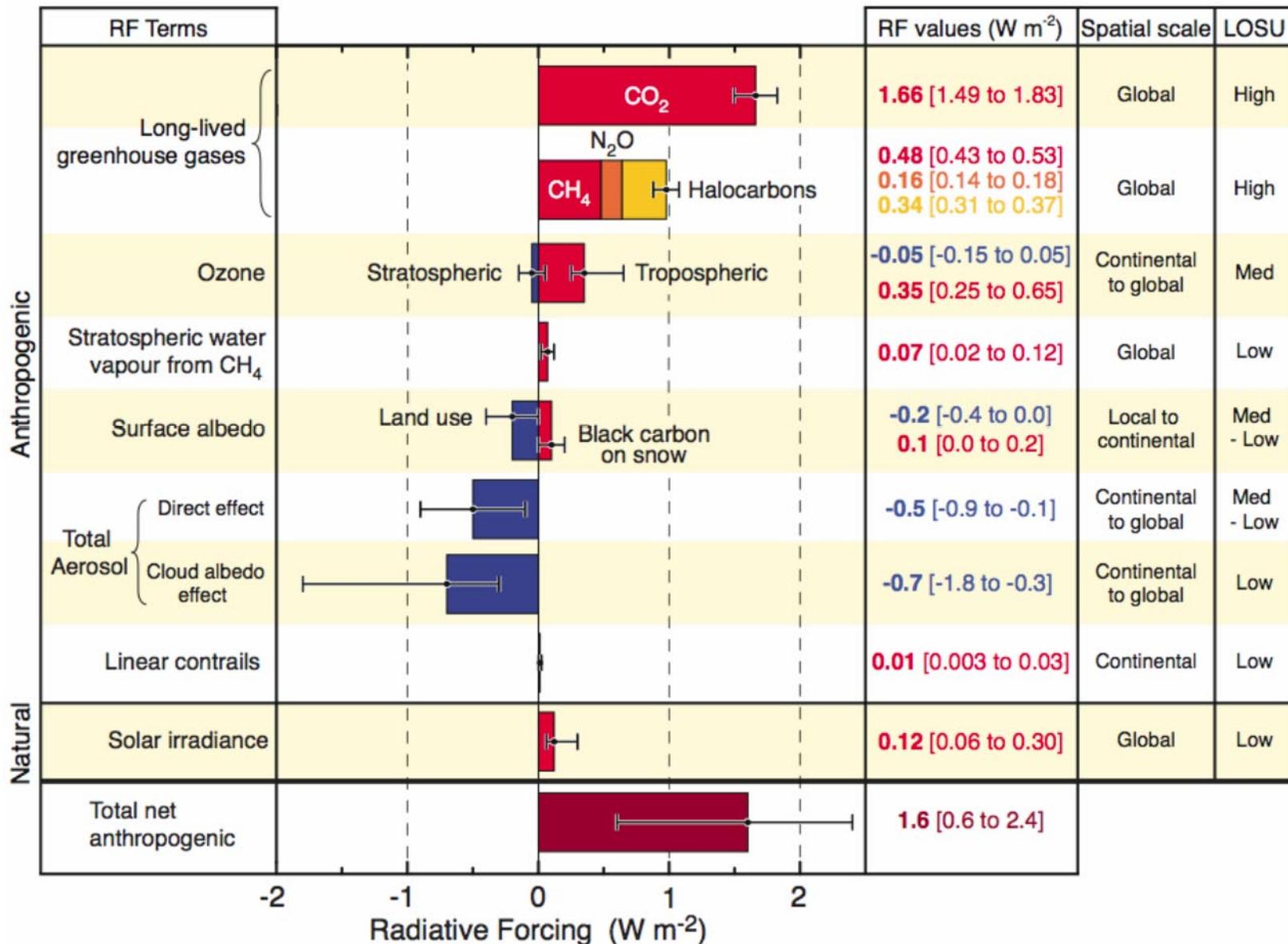
*Climate sensitivity from paleo climate has been a major contributor to present assessment of climate sensitivity.*

# ***KEY APPROACHES TO DETERMINING CLIMATE SENSITIVITY***

- ***Paleoclimate studies***: Forcing and response over time scales from millennial to millions of years.
- ***Empirical***: Forcing and response over the instrumental record.

# GLOBAL-MEAN RADIATIVE FORCINGS (RF)

Pre-industrial to present (Intergovernmental Panel on Climate Change, 2007)

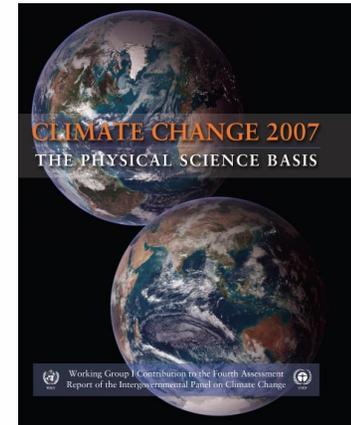
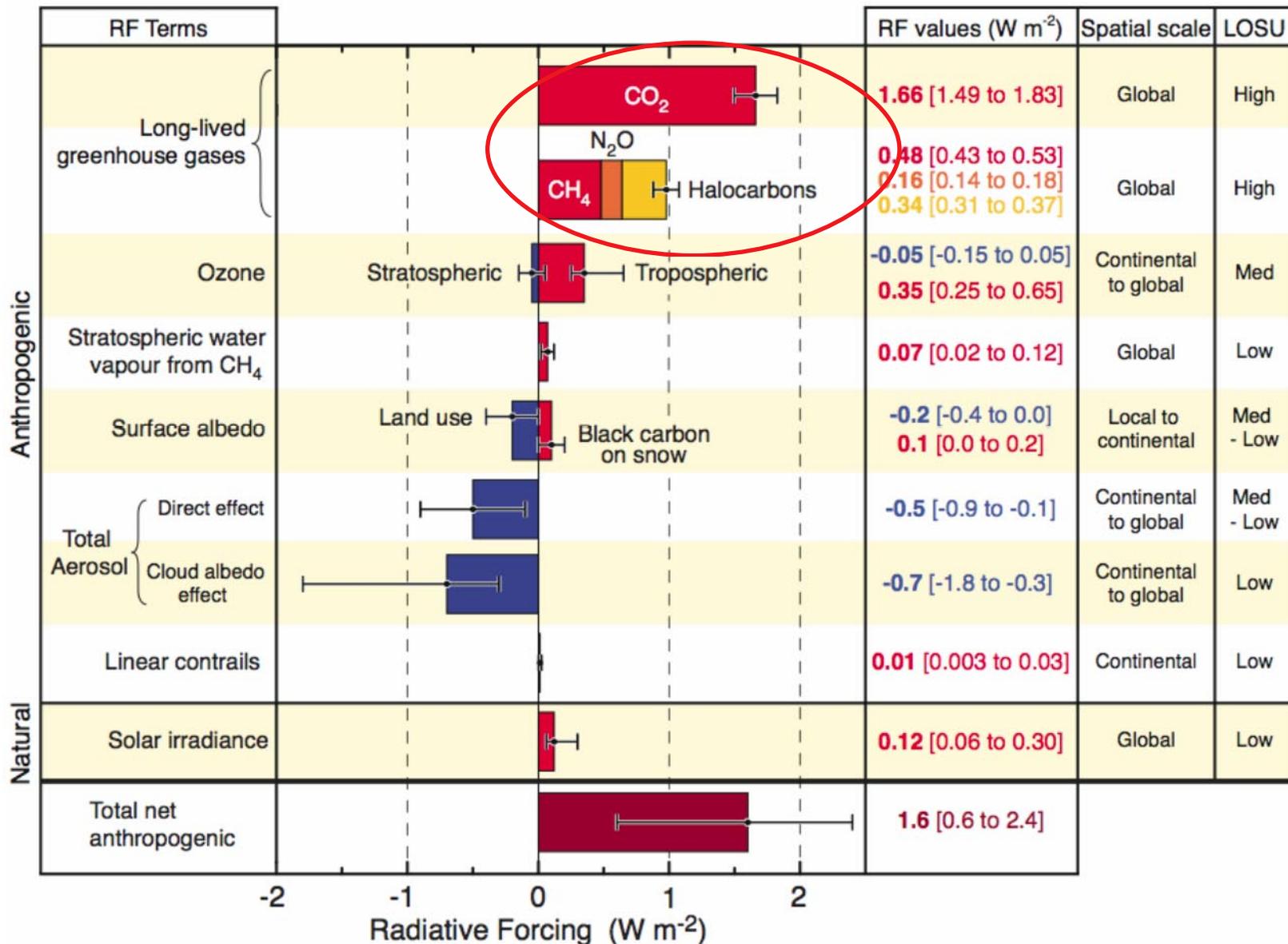


©IPCC 2007: WG1-AR4

LOSU denotes level of scientific understanding.

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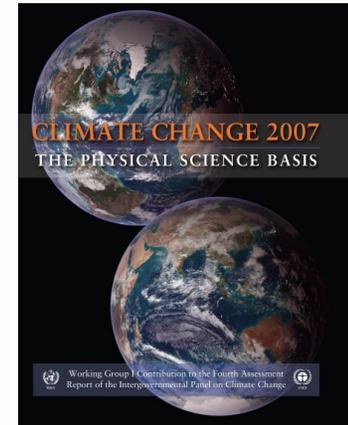
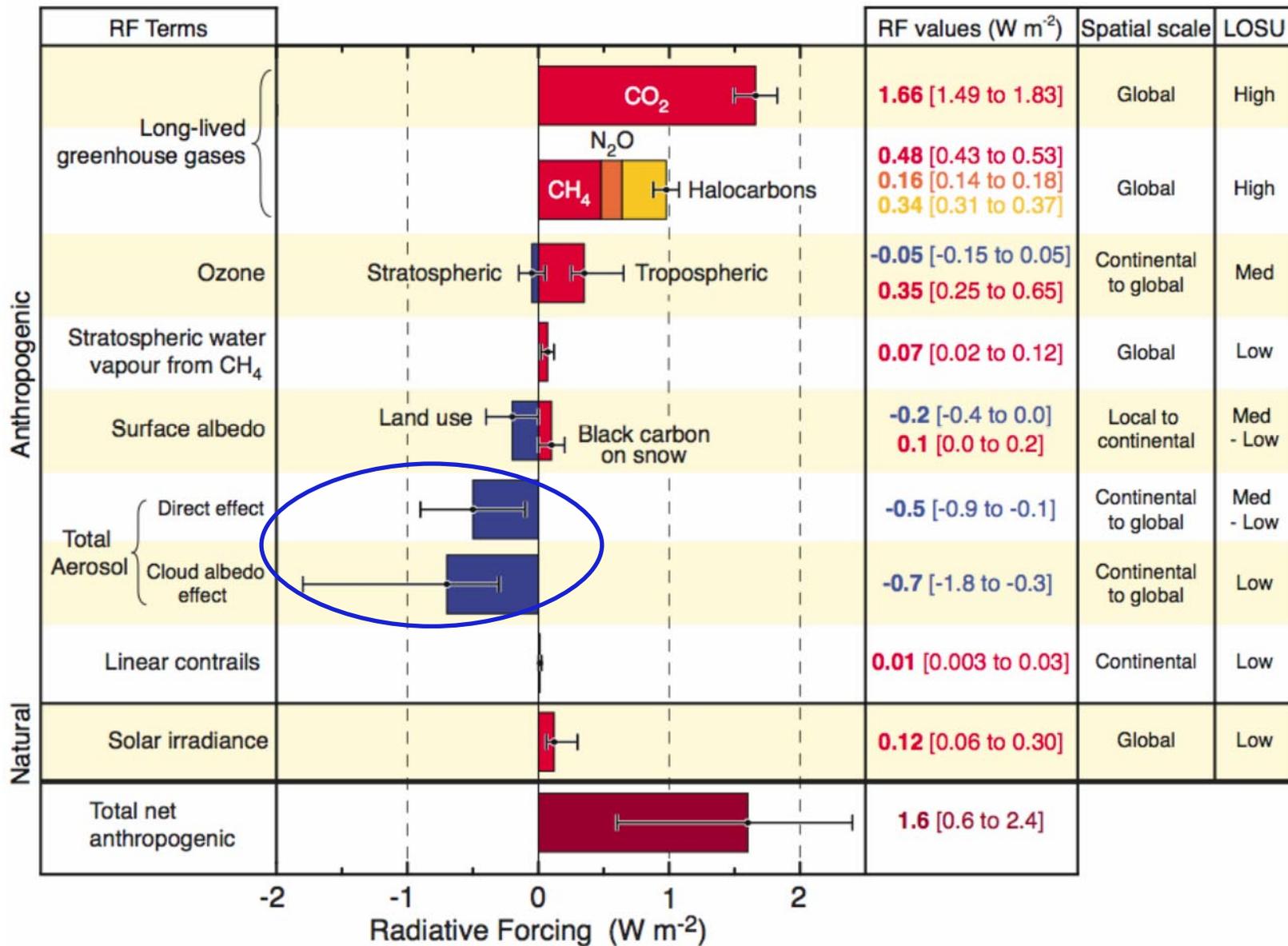


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# GLOBAL-MEAN RADIATIVE FORCINGS (RF)

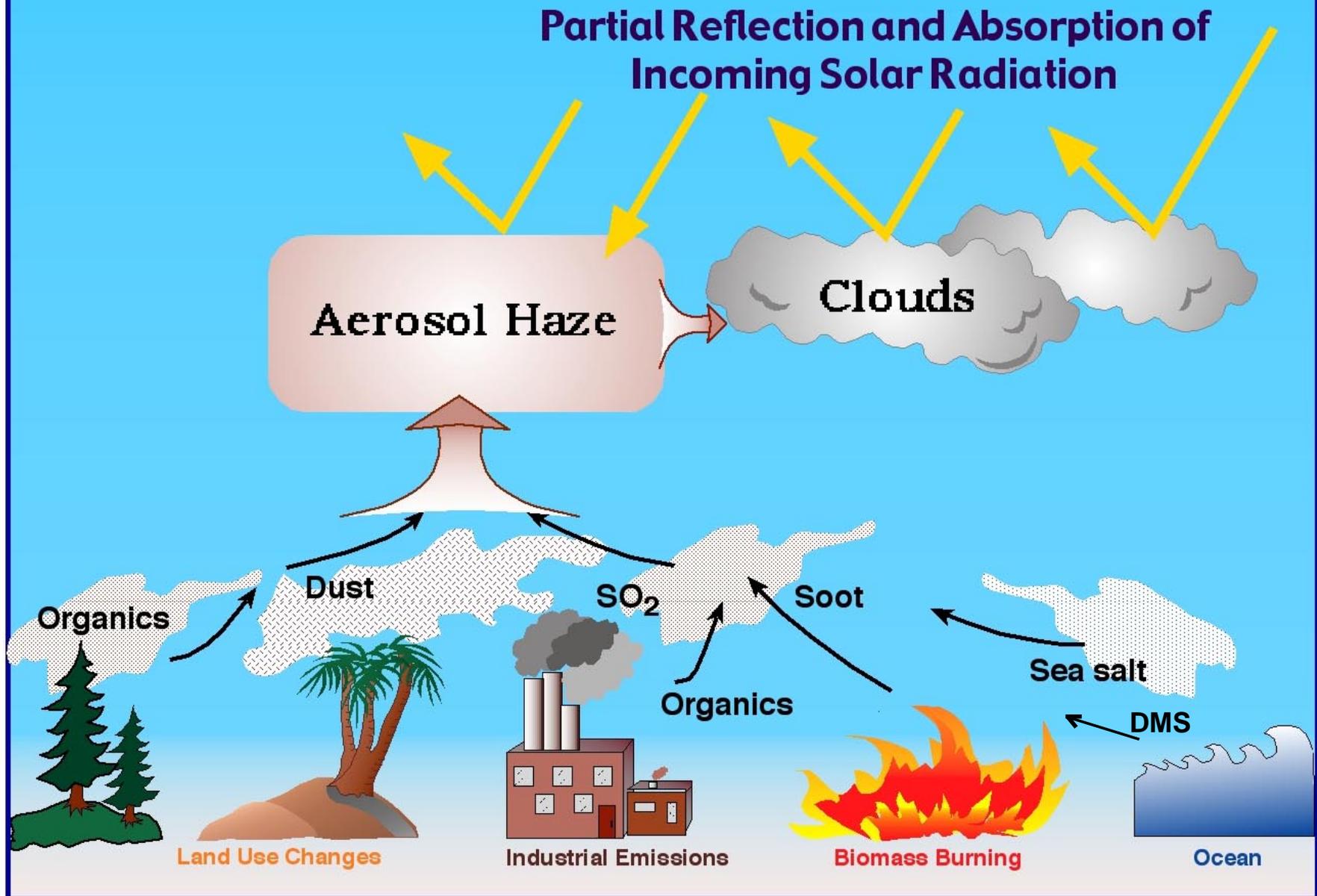
Pre-industrial to present (Intergovernmental Panel on Climate Change, 2007)



©IPCC 2007: WG1-AR4

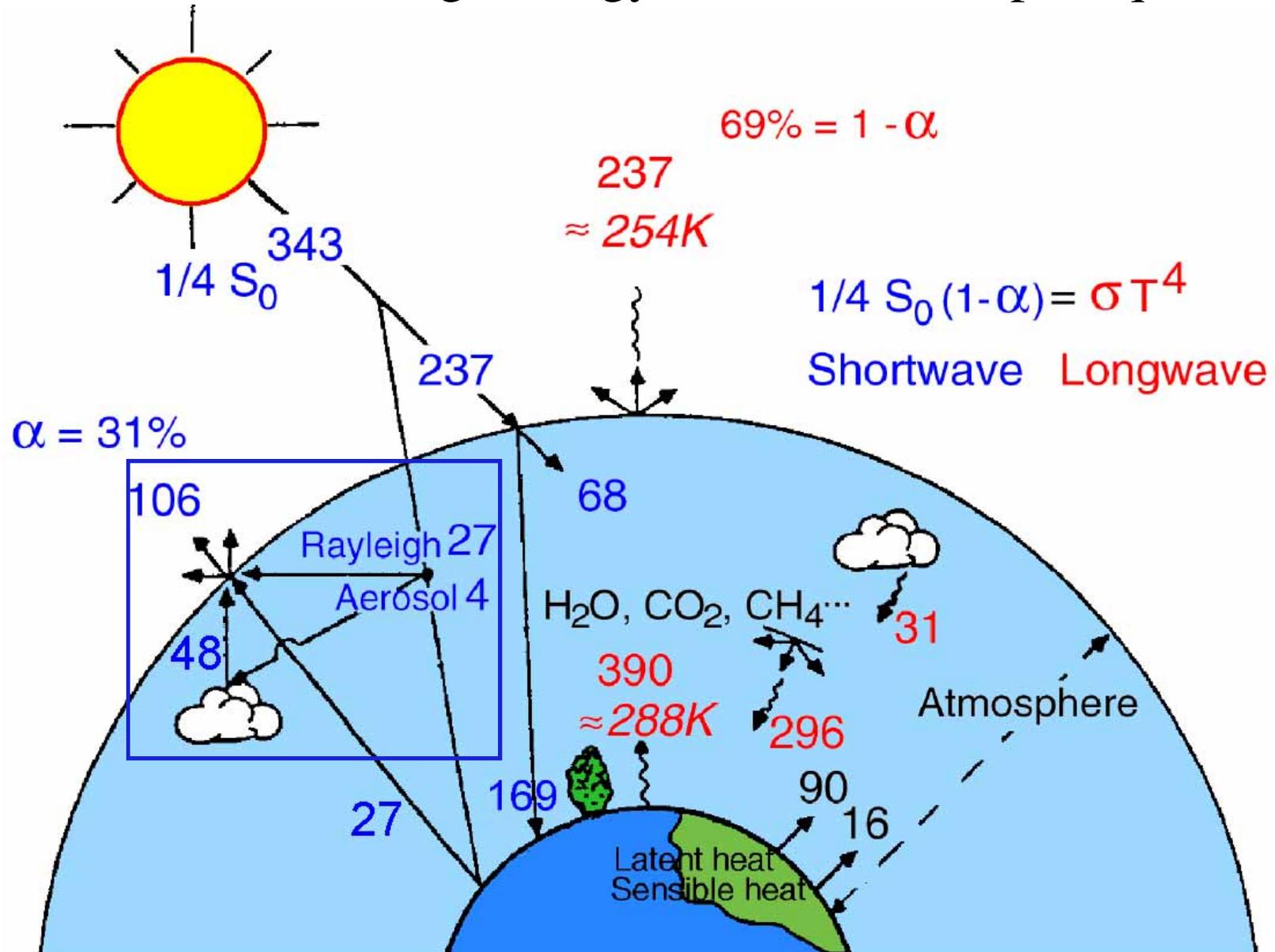
LOSU denotes level of scientific understanding.

# Radiative Forcing by Tropospheric Aerosol



# GLOBAL ENERGY BALANCE

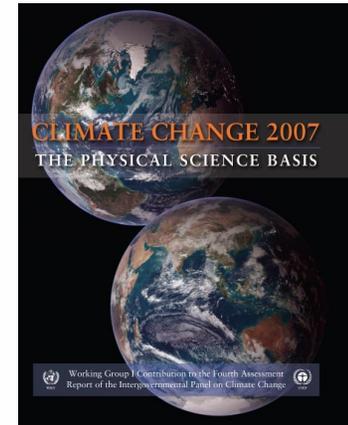
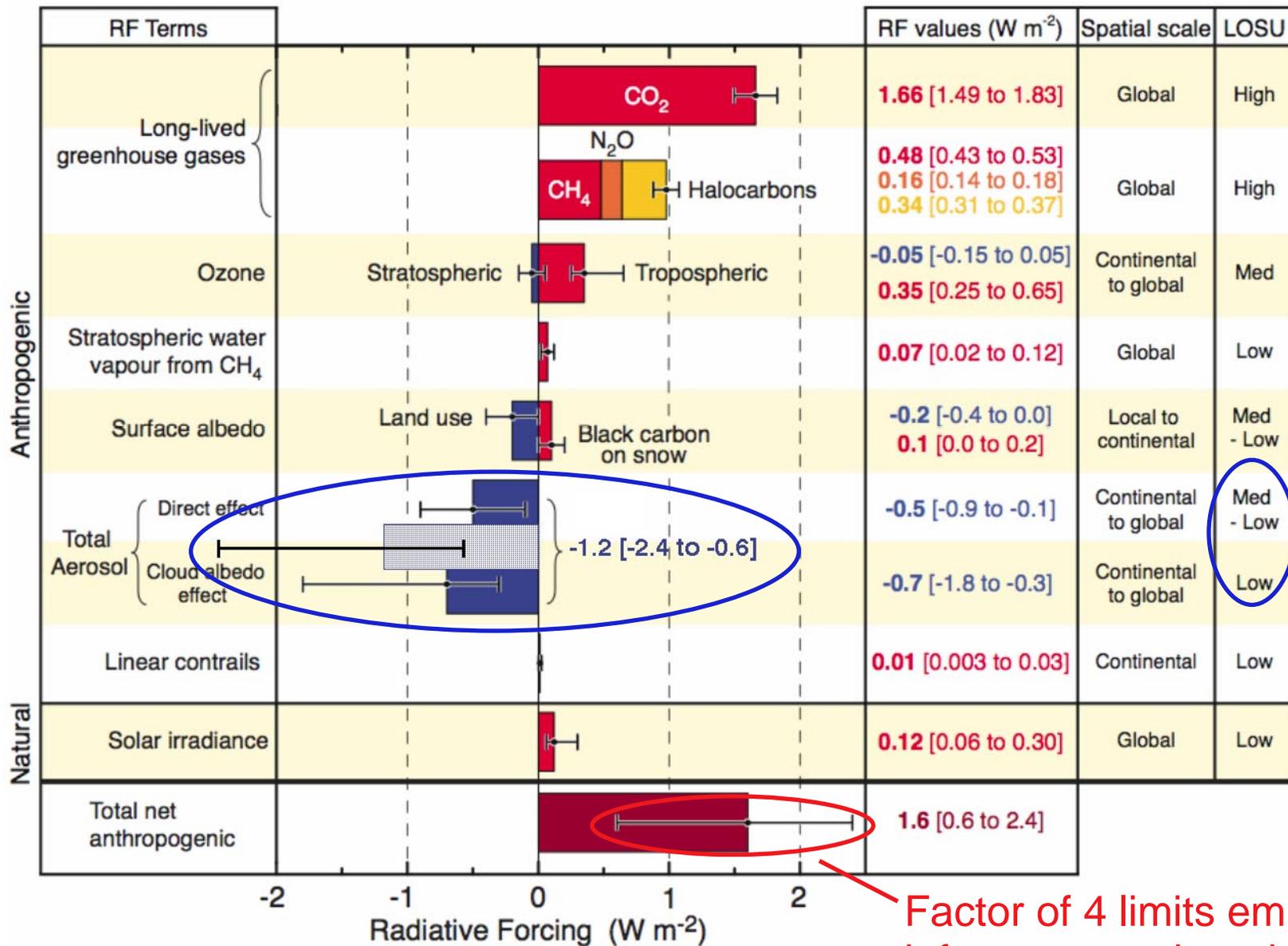
Global and annual average energy fluxes in watts per square meter



*Schwartz, 1996, modified from Ramanathan, 1987*

# GLOBAL-MEAN RADIATIVE FORCINGS (RF)

Pre-industrial to present (Intergovernmental Panel on Climate Change, 2007)



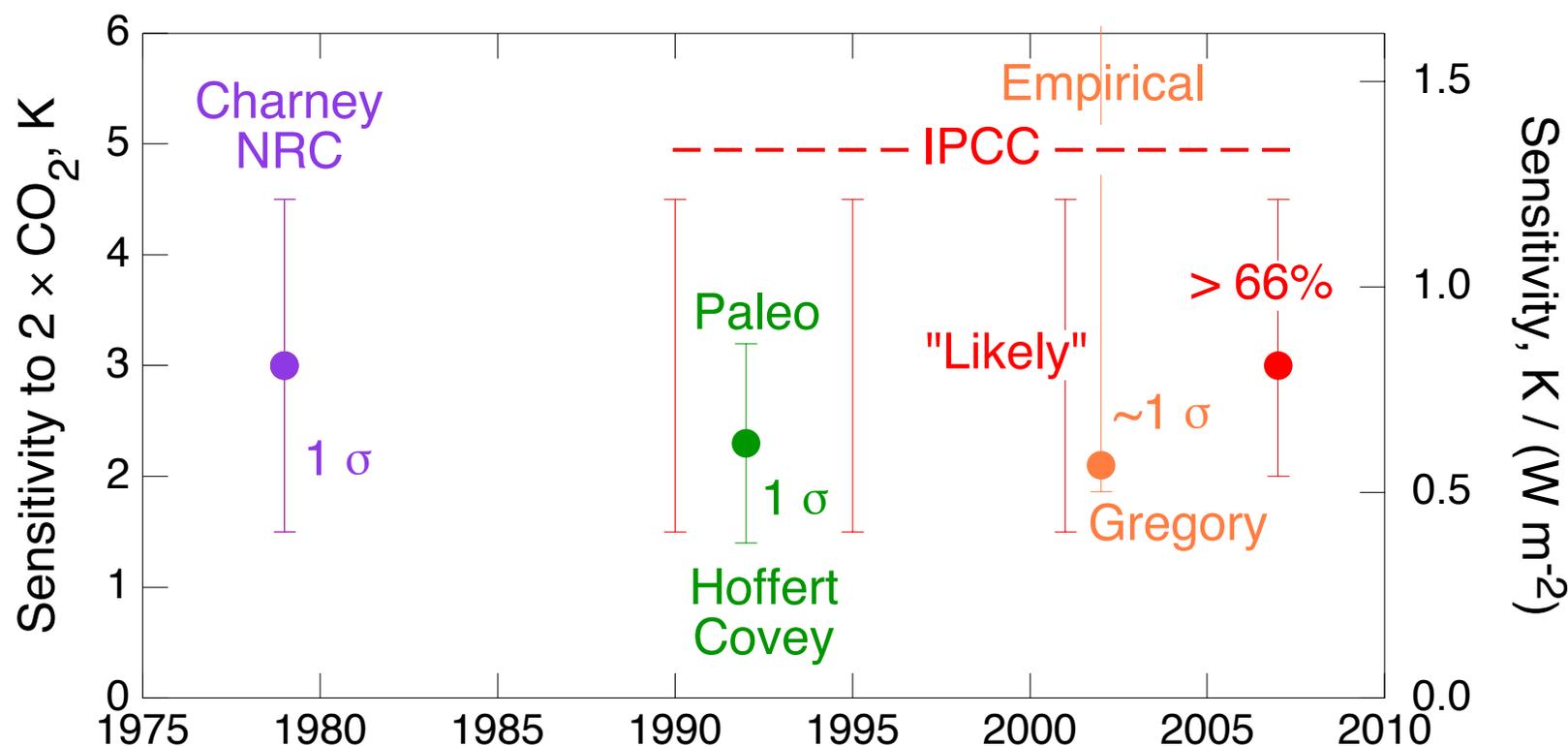
©IPCC 2007: WG1-AR4

Factor of 4 limits empirical inferences and model evaluation.

LOSU denotes level of scientific understanding.

# CLIMATE SENSITIVITY ESTIMATES THROUGH THE AGES

Estimates of central value and uncertainty range from specific approaches and major national and international assessments



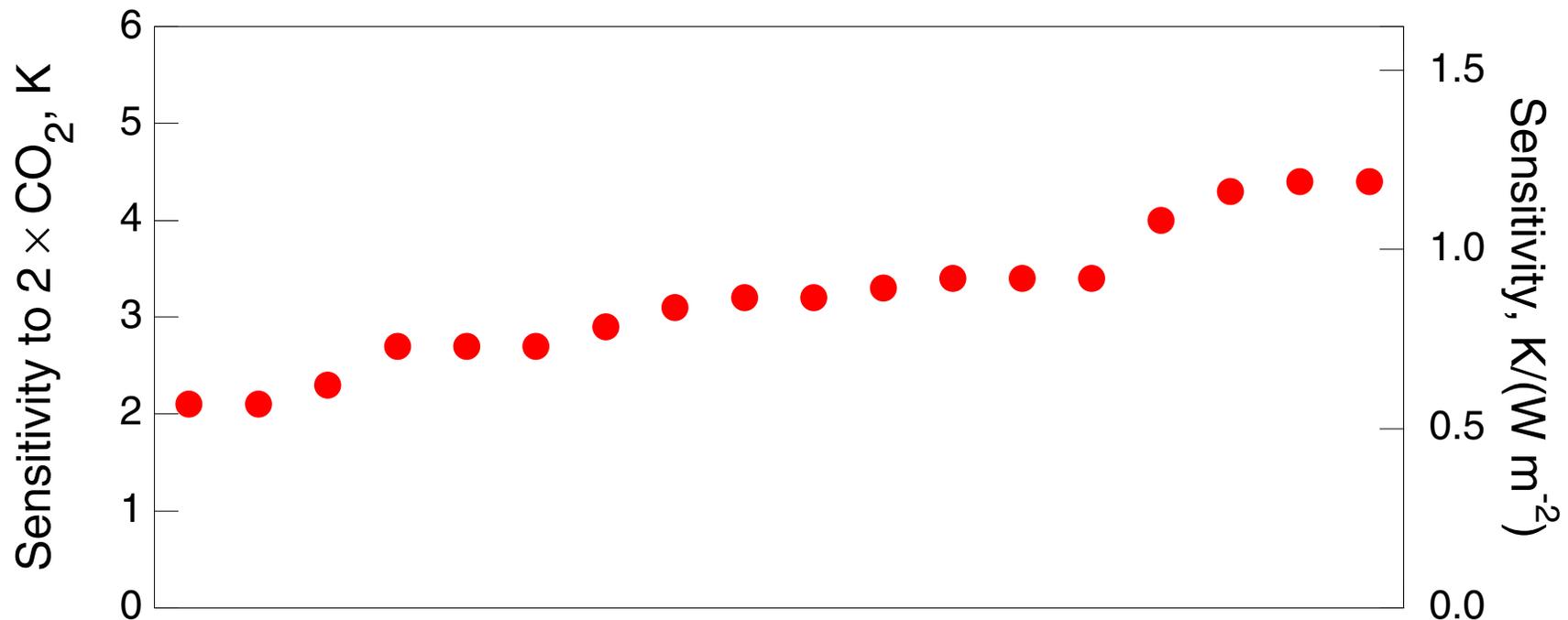
*Empirical approach does not greatly constrain sensitivity because of uncertainty in aerosol forcing over the period of instrumental record.*

# ***KEY APPROACHES TO DETERMINING CLIMATE SENSITIVITY***

- ***Paleoclimate studies***: Forcing and response over time scales from millennial to millions of years.
- ***Empirical***: Forcing and response over the instrumental record.
- ***Climate modeling***: Understanding the processes that comprise Earth's climate system and representing them in large-scale numerical models.

# CLIMATE SENSITIVITY ESTIMATES FROM GLOBAL CLIMATE MODELS

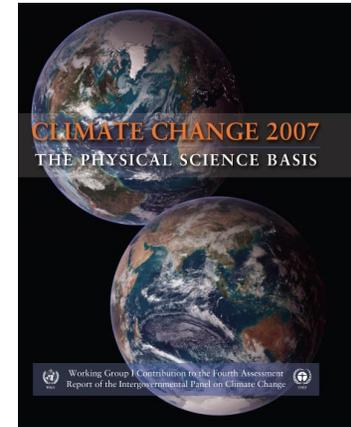
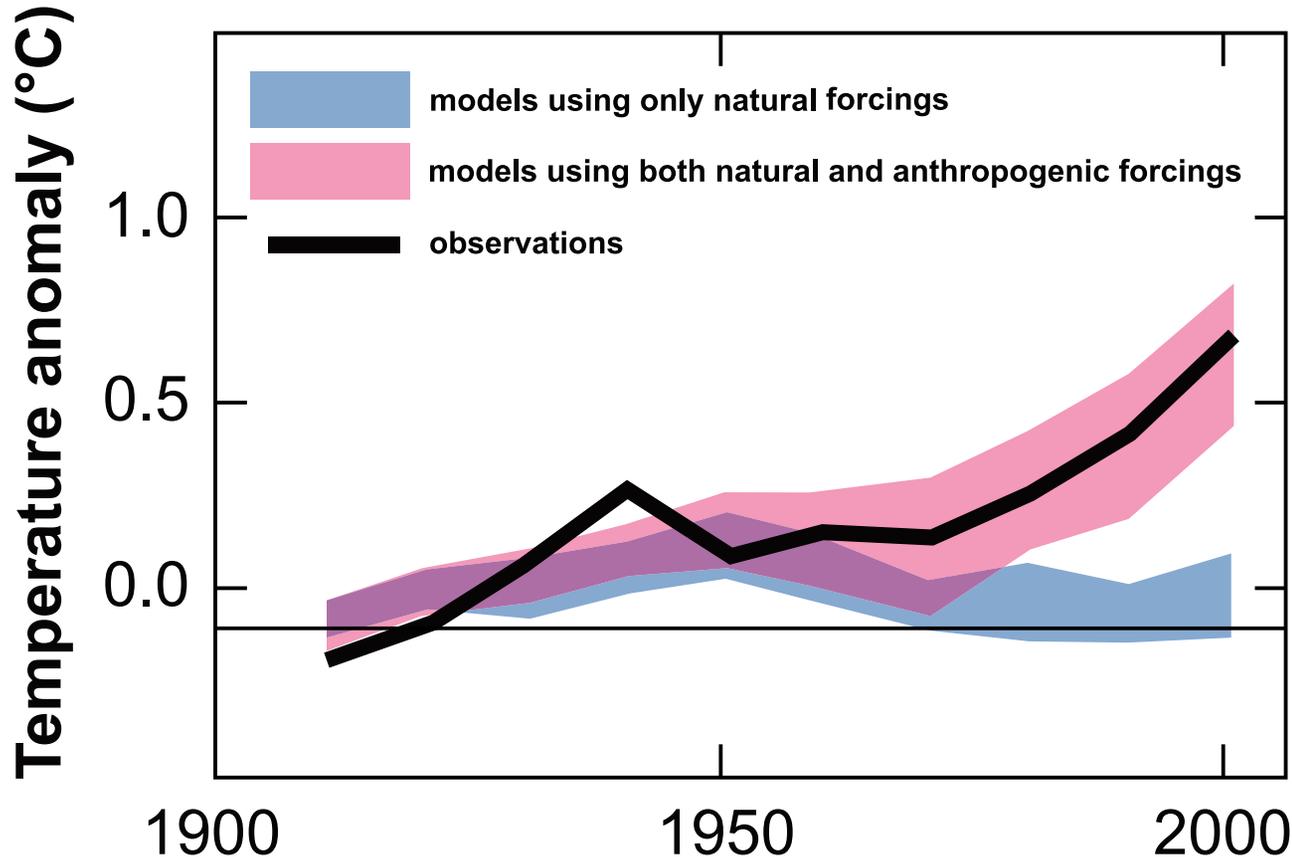
18 Current global climate models – IPCC AR4, 2007



*Range of model sensitivities is identical with range of current overall IPCC sensitivity estimate.*

# TOO ROSY A PICTURE?

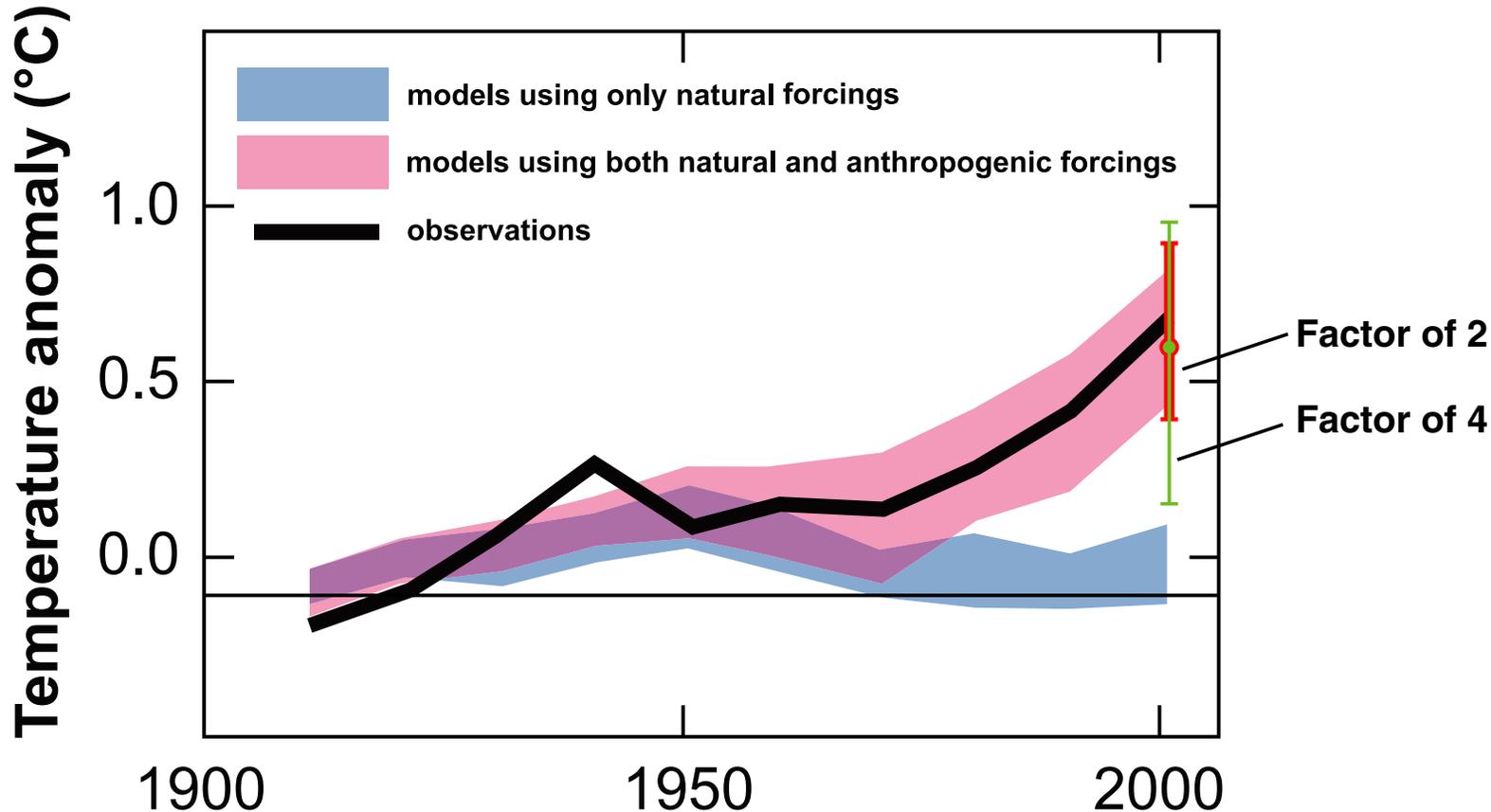
Ensemble of 58 model runs with 14 global climate models



- “ Simulations that incorporate anthropogenic forcings, including increasing greenhouse gas concentrations and the effects of aerosols, and that also incorporate natural external forcings provide a *consistent explanation of the observed temperature record*.
- “ These simulations used models with *different climate sensitivities, rates of ocean heat uptake and magnitudes and types of forcings*.

# TOO ROSY A PICTURE?

Ensemble of 58 model runs with 14 global climate models



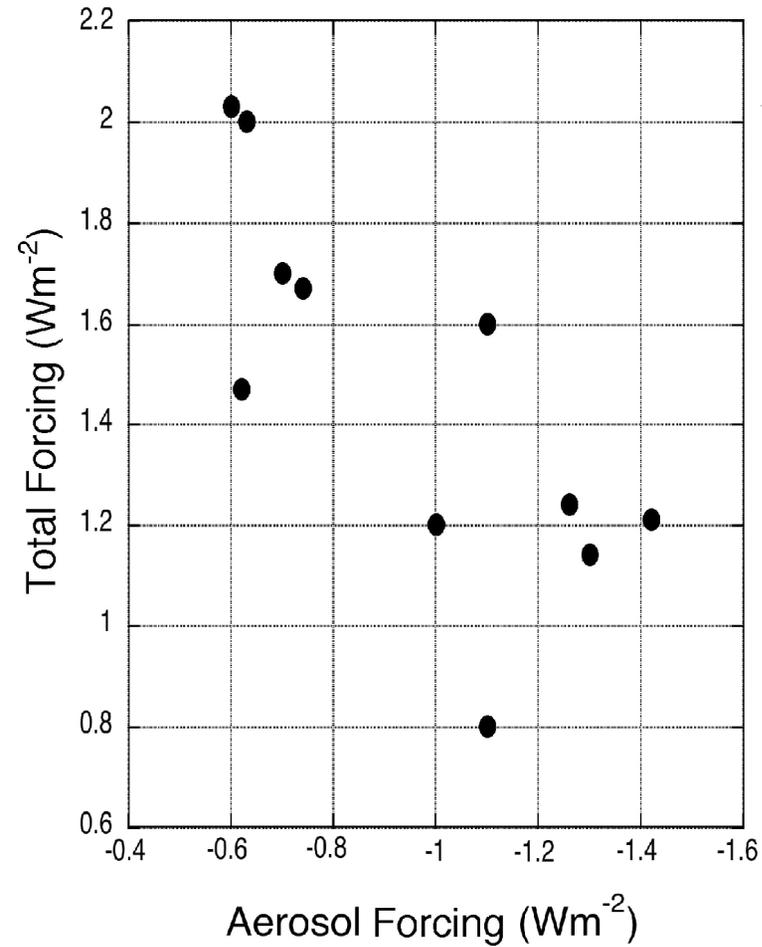
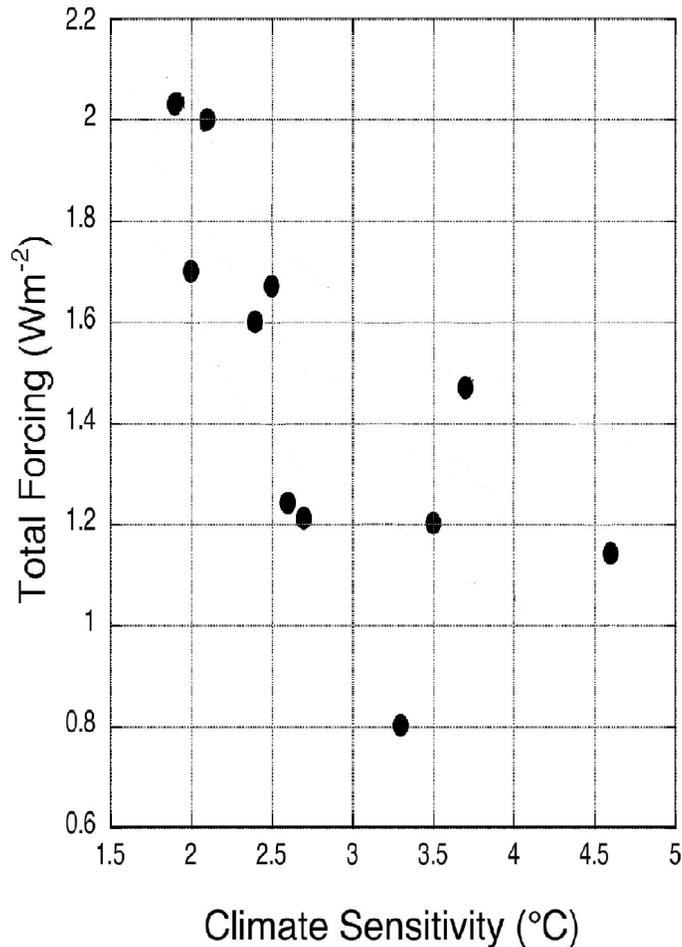
*Schwartz, Charlson & Rodhe, Nature Reports – Climate Change, 2007*

The models *did not span the full range of the uncertainty* and/or . . .

The forcings used in the model runs were *anticorrelated with the sensitivities of the models*.

# CORRELATION OF AEROSOL FORCING, TOTAL FORCING, AND SENSITIVITY IN CLIMATE MODELS

Eleven models used in 2007 IPCC analysis



*Modified from Kiehl, GRL, 2007*

Climate models with higher sensitivity have lower total forcing.  
Lower total forcing is due to greater (negative) aerosol forcing.

# ***KEY APPROACHES TO DETERMINING CLIMATE SENSITIVITY***

- ***Paleoclimate studies***: Forcing and response over time scales from millennial to millions of years.
- ***Empirical***: Forcing and response over the instrumental record.
- ***Climate modeling***: Understanding the processes that comprise Earth's climate system and representing them in large-scale numerical models.
- ***Energy-balance model***: Empirical determination from integral properties of Earth's climate system.

# ENERGY BALANCE MODEL OF EARTH'S CLIMATE SYSTEM



$$\text{Global energy balance: } C \frac{dT_s}{dt} = \frac{dH}{dt} = Q - E = \frac{\gamma J_S}{4} - \epsilon \sigma T_s^4$$

$C$  is heat capacity coupled to climate system on relevant time scale

$T_s$  is global mean surface temperature

$H$  is global heat content

$Q$  is absorbed solar energy

$E$  is emitted longwave flux

$J_S$  is solar constant

$\gamma$  is planetary co-albedo

$\sigma$  is Stefan-Boltzmann constant

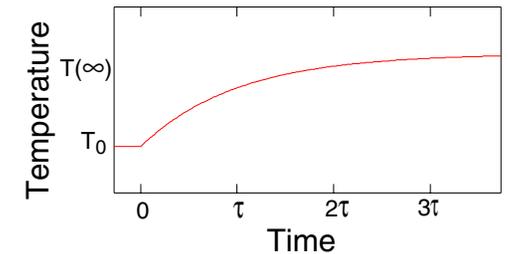
$\epsilon$  is effective emissivity

# ENERGY BALANCE MODEL OF EARTH'S CLIMATE SYSTEM



Apply step-function forcing:  $\Delta F = \Delta(Q - E)$

At new “equilibrium”  $\Delta T_s(\infty) = S\Delta F$



$S$  is equilibrium climate sensitivity

$$S = \frac{T_0}{\gamma_0 J_S} \frac{1}{\left(1 - \frac{1}{4} \frac{d \ln \gamma}{d \ln T} \Big|_0 + \frac{1}{4} \frac{d \ln \varepsilon}{d \ln T} \Big|_0\right)} \quad \text{K / (W m}^{-2}\text{)}$$

If  $\gamma$  and  $\varepsilon$  are constant (*no feedbacks*),  $S = \frac{T_0}{\gamma_0 J_S}$  Stefan-Boltzmann sensitivity,  $S_{SB} = 0.30 \text{ K / (W m}^{-2}\text{)}$ ;  $\Delta T_{2\times} = 1.1 \text{ K}$

$f$  is feedback factor

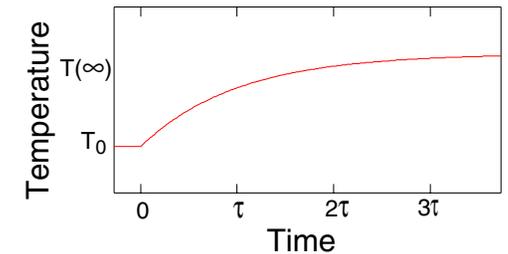
$$f = \frac{1}{\left(1 - \frac{1}{4} \frac{d \ln \gamma}{d \ln T} \Big|_0 + \frac{1}{4} \frac{d \ln \varepsilon}{d \ln T} \Big|_0\right)}$$

# ENERGY BALANCE MODEL OF EARTH'S CLIMATE SYSTEM



Apply step-function forcing:  $\Delta F = \Delta(Q - E)$

At “equilibrium”  $\Delta T_s(\infty) = S\Delta F$



$S$  is equilibrium climate sensitivity  $S = \frac{T_0}{\gamma_0 J_S} f = S_{\text{SB}} f$  Stefan-Boltzmann sensitivity times feedback factor

*Time dependence:*  $\Delta T_s(t) = S\Delta F(1 - e^{-t/\tau})$

$\tau$  is climate system time constant  $\tau = CS$  or  $S = \tau / C$

***One equation in three unknowns!***

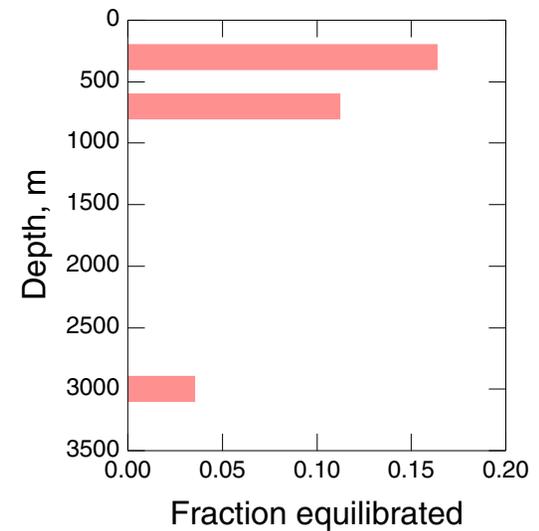
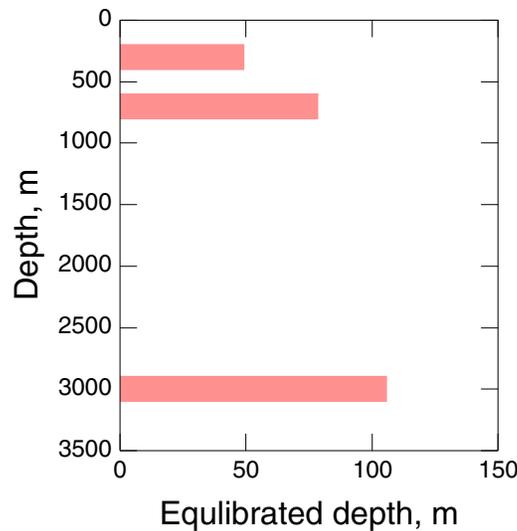
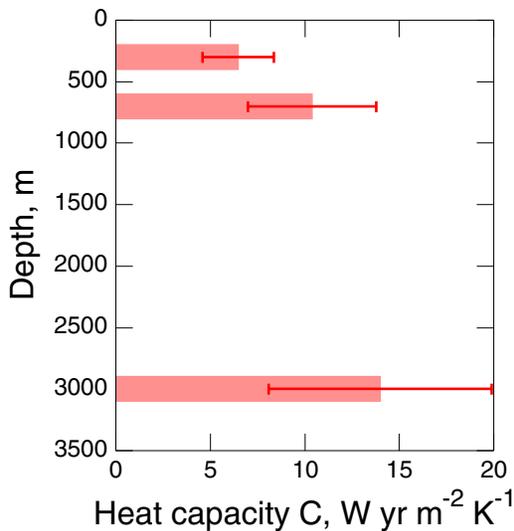
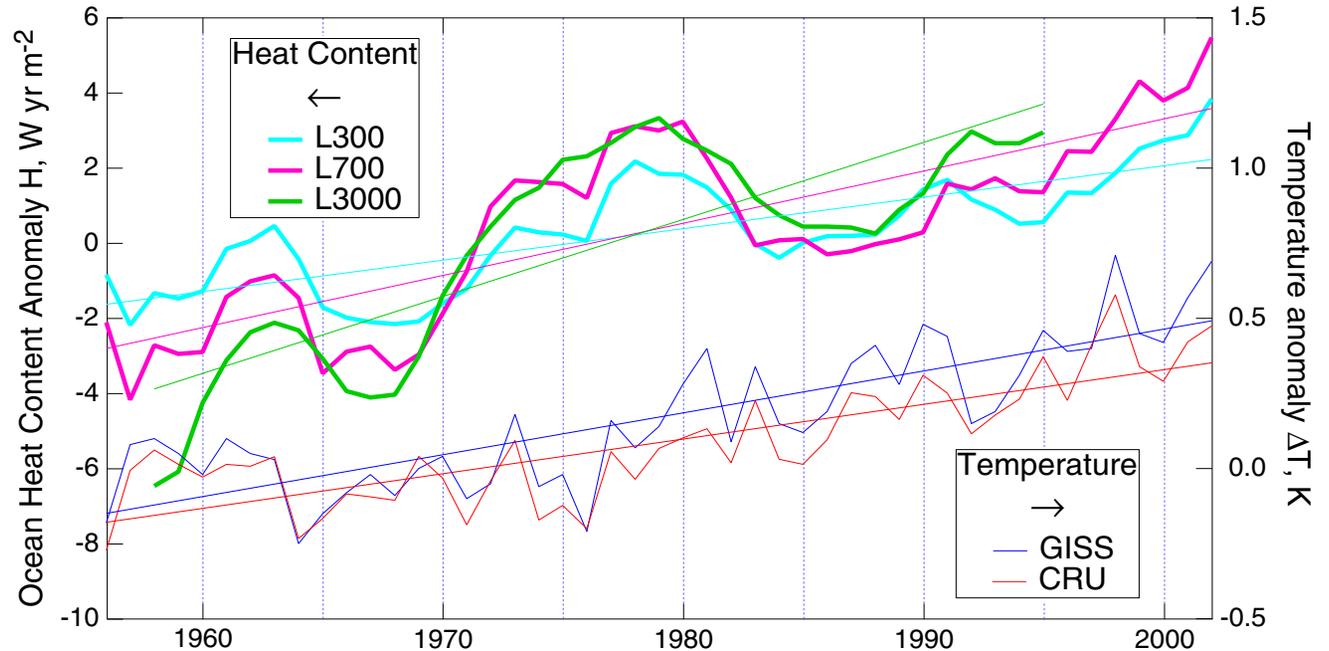
***Approach:*** Determine  $C$  and  $\tau$  from measurements; calculate sensitivity  $S$ .

# EMPIRICAL DETERMINATION OF OCEAN HEAT CAPACITY

$$C = \frac{dH / dt}{dT_s / dt}$$

Ocean heat content  
*H*: Levitus *et al.*,  
 GRL, 2005

Surface temperature  
*T<sub>s</sub>*: GISS, CRU

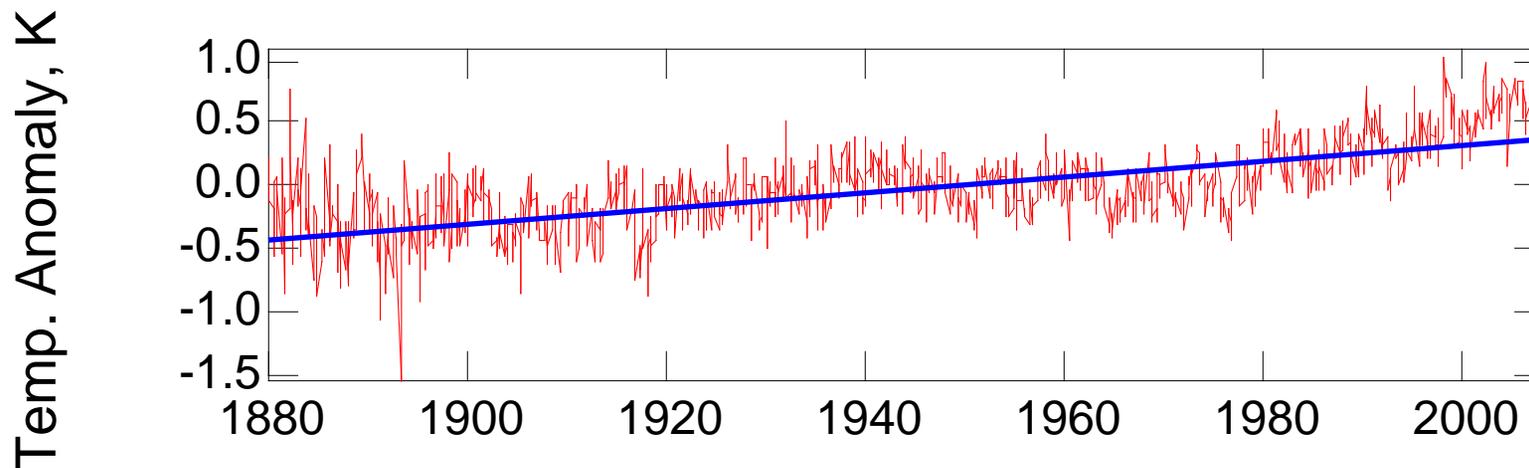


- ~50% of heat capacity is between surface and 300 m.
- Other heat sinks raise global heat capacity to  $17 \pm 7 \text{ W yr m}^{-2} \text{ K}^{-1}$ .

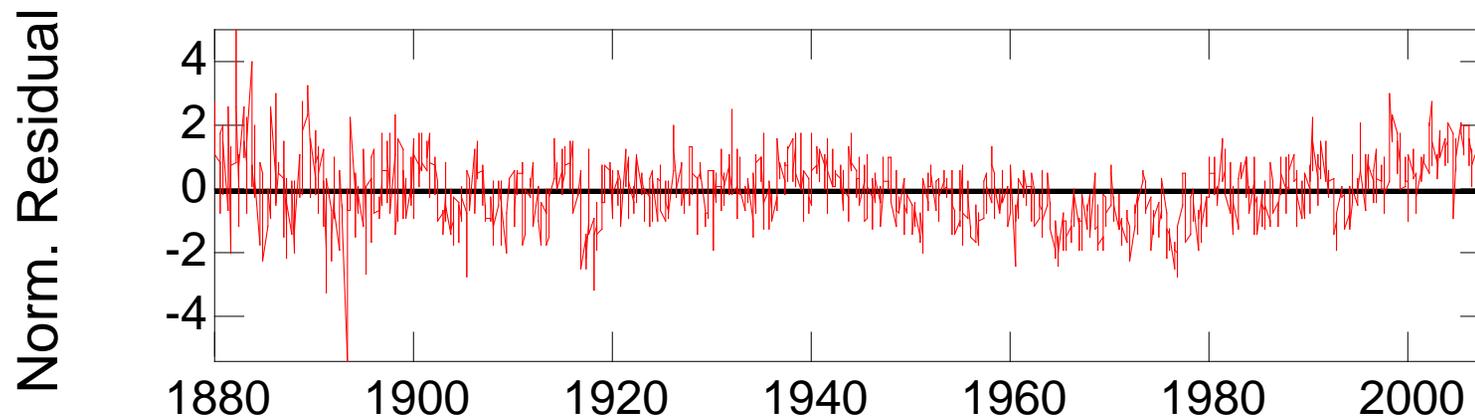
# TIME CONSTANT OF EARTH'S CLIMATE SYSTEM

Determination from autocorrelation of time series

*Input:* Monthly global-mean surface temperature anomaly  $T_s$



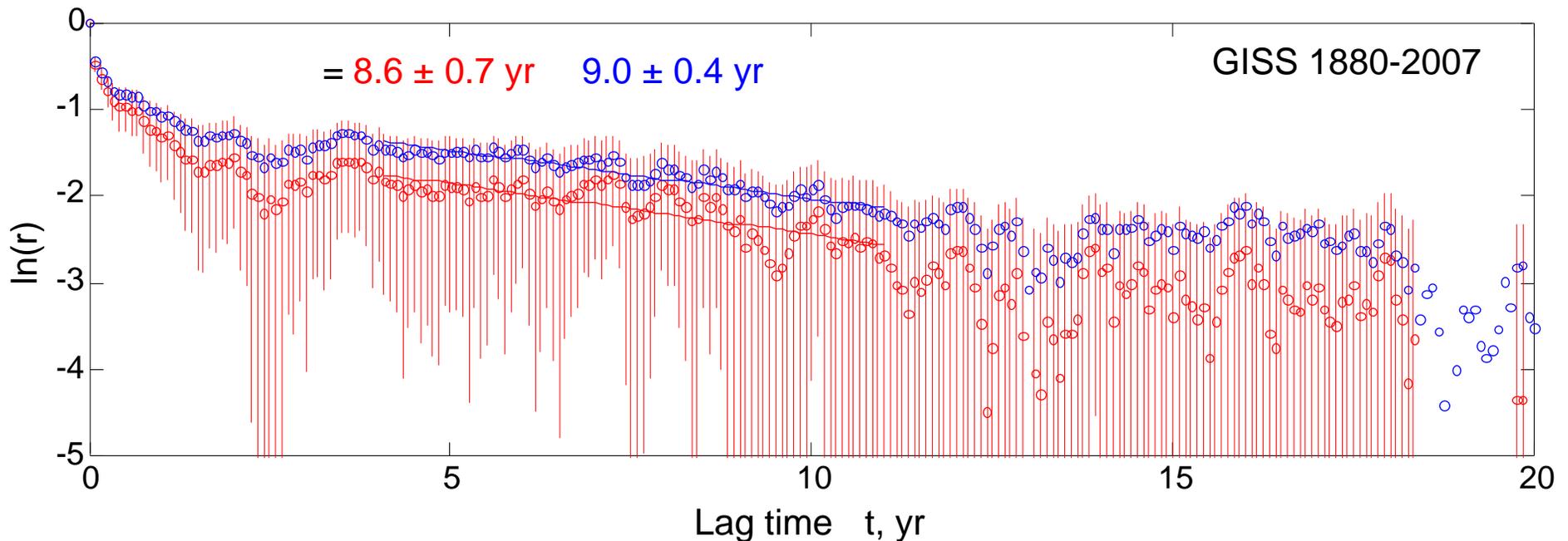
Remove long term trend; plot the residuals:



# TIME CONSTANT OF EARTH'S CLIMATE SYSTEM

Determination from autocorrelation of time series (*cont'd*)

Evaluate *climate system time constant* as  $\tau = (d \ln r(t) / d t)^{-1}$   
*Correct for short duration of time series.*



Summary (multiple data sets):

Climate system time constant is  $8.5 \pm 2.5 \text{ years}$

# EVALUATION OF SENSITIVITY AND FORCINGS

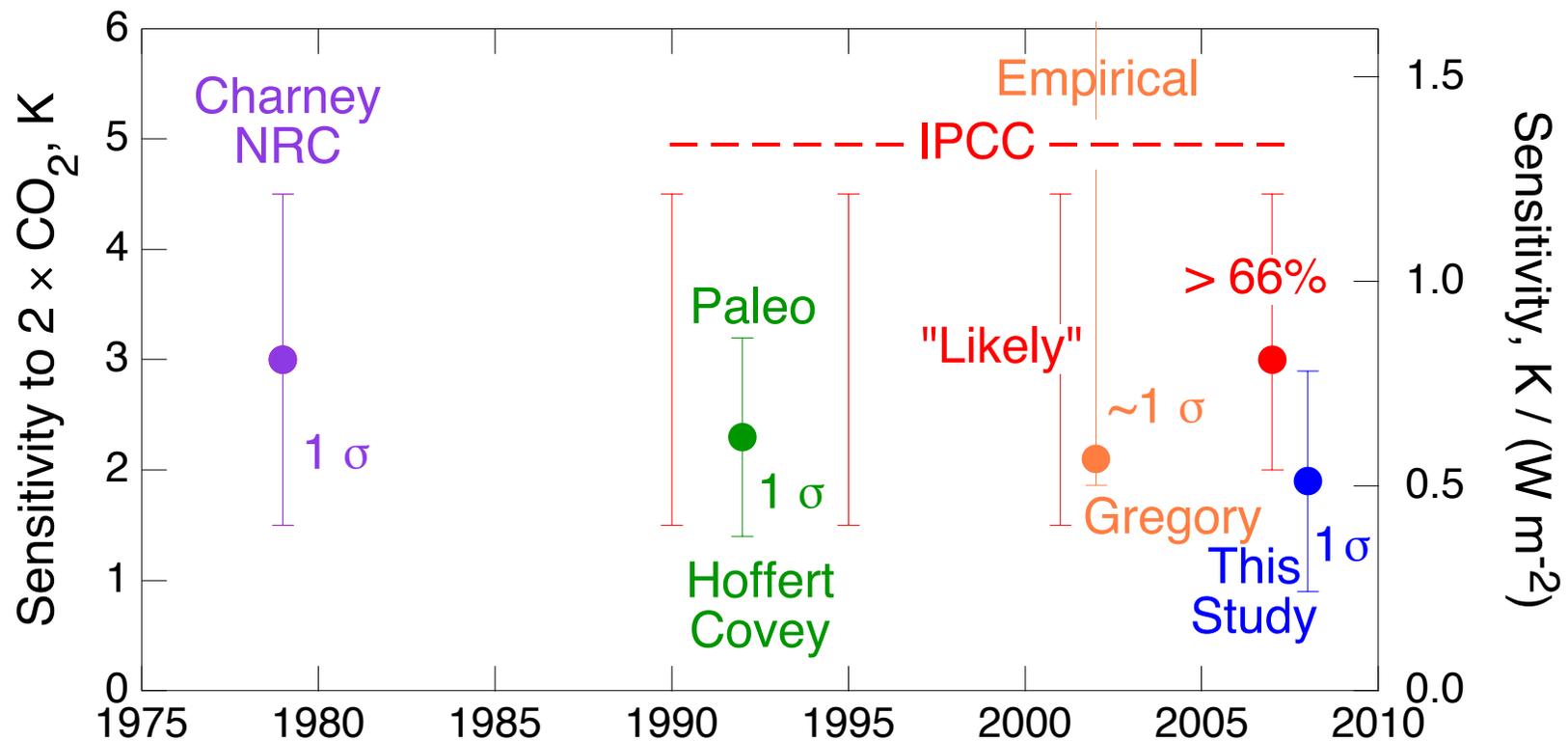
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Quantity	Unit	Value	1 $\sigma$
Effective global heat capacity $C$	W yr m <sup>-2</sup> K <sup>-1</sup>	<b>17</b>	<b>7</b>
Effective climate system time constant $\tau$	yr	<b>8.5</b>	<b>2.5</b>
Equilibrium climate sensitivity $S = \tau / C$	K/(W m <sup>-2</sup> )	<b>0.51</b>	<b>0.26</b>
Feedback factor $f$	–	<b>1.7</b>	
Equilibrium temperature increase for $2 \times \text{CO}_2$ , $\Delta T_{2\times}$	K	<b>1.9</b>	<b>1.0</b>

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# CLIMATE SENSITIVITY ESTIMATES THROUGH THE AGES

Estimates of central value and uncertainty range from specific approaches and major national and international assessments



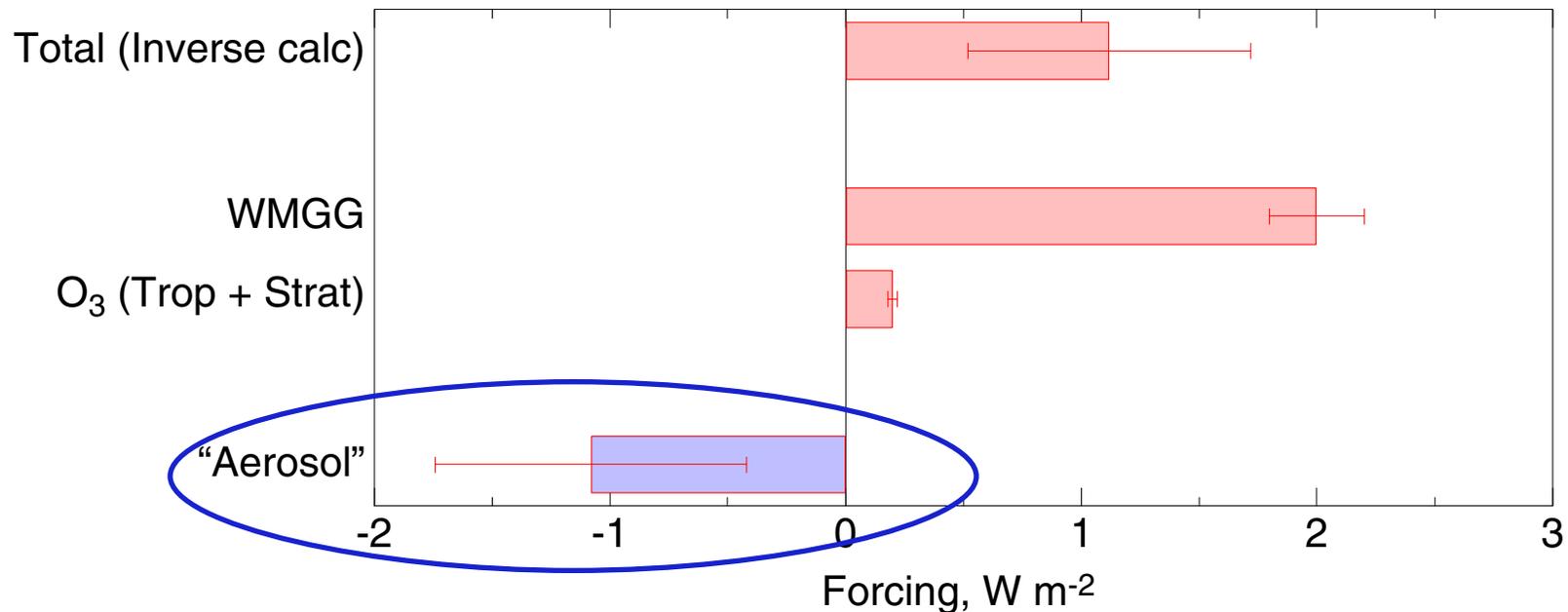
*Sensitivity obtained in this study overlaps range from climate models, paleo, empirical; seems to rule out  $\Delta T_{2x} \gtrsim 3$  K.*

# EVALUATION OF SENSITIVITY AND FORCINGS

Quantity	Unit	Value	1 $\sigma$
Effective global heat capacity $C$	$\text{W yr m}^{-2} \text{K}^{-1}$	<b>17</b>	<b>7</b>
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Feedback factor $f$	–	<b>1.7</b>	
Equilibrium temperature increase for $2 \times \text{CO}_2$ , $\Delta T_{2\times}$	K	<b>1.9</b>	<b>1.0</b>
Total forcing over the 20 <sup>th</sup> century, $F_{20} = \Delta T_{20} / S$ <i>Inverse calculation</i>	$\text{W m}^{-2}$	<b>1.1</b>	<b>0.6</b>
Forcing in 20 <sup>th</sup> century other than GHGs (mainly aerosols), $F_{20}^{\text{other}} = F_{20} - F_{20}^{\text{ghg}}$	$\text{W m}^{-2}$	<b>-1.1</b>	<b>0.7</b>

# INVERSE CALCULATION OF “AEROSOL” FORCING OVER TWENTIETH CENTURY

$$\text{“Aerosol” forcing} = \text{Total forcing} - \text{GHG forcing}$$



Total forcing remains uncertain to a factor of 3.

“Aerosol” forcing, calculated as residual, is presumably dominated by aerosols.

“Aerosol” forcing is substantial, with large uncertainty.

“Aerosol” forcing could be masking as much as 75% of GHG warming.

# EVALUATION OF SENSITIVITY AND FORCINGS

Quantity	Unit	Value	1 $\sigma$
Effective global heat capacity $C$	W yr m <sup>-2</sup> K <sup>-1</sup>	<b>17</b>	<b>7</b>
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Forcing in 20 <sup>th</sup> century other than GHGs (mainly aerosols), $F_{20}^{\text{other}} = F_{20} - F_{20}^{\text{ghg}}$	W m <sup>-2</sup>	<b>-1.1</b>	<b>0.7</b>
Lag in temperature change, $\Delta T_{\text{lag}}$	K	<b>0.05</b>	

# CONCLUDING REMARKS

Traditional approaches to determination of Earth's climate sensitivity yield uncertainty of at least a factor of 3, largely because of uncertainty in aerosol forcing.

The energy balance approach offers a new independent determination of Earth's climate sensitivity that does not depend on knowledge of aerosol forcing.

This approach yields a sensitivity that is at the low end of current estimates and would seem to rule out high sensitivity.

The short time constant,  $\sim 8.5$  years, suggests little heating in the pipeline from time lags.

Aerosols could be masking up to 75% of GHG forcing and warming.

Nothing in the present study should be construed as diminishing the need for strenuous reduction in GHG emissions.

# FINAL REMARKS

This study is a first effort on this approach. I would hope that it would be refined by further research.

Would I bet the ranch on this analysis? Of course not.